

RESEARCH ARTICLE

Hydrogeochemical and Groundwater Assessment for Drinking Purpose at ITERA Campus Area and Its Surroundings

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Abstract

The total population around ITERA has increased every year as students acceptance every year. To anticipate this, it needs to be done a research at ITERA campus and its surrounding about the quality of groundwater for drinking purpose and the hydrogeochemical of groundwater to know the controlling factors which are dominant. The methods are integrating Piper diagram plotting result, X-Y plotting result for some cations and anions, and Gibbs diagram plotting result. It is for hydrogeochemical analysis. Groundwater assessment for drinking purpose referred to Peraturan Menteri Kesehatan Republik Indonesia No. 492/MENKES/PER/IV/2010. There were 14 samples that were taken from nine dig wells and five drill wells. The groundwater facieses were dominated by Facies Na-HCO₃-Cl (35,71%) followed by Facies Na-Cl (21,43%), Facies Na-HCO₃ (21,43%), Facies Na-SO₄-Cl (14,29%), dan Facies Ca-Mg-HCO₃ (7,14%). Groundwater hydrogeochemical of research area shows that groundwater chemistries are controlled by minerals weathering, evaporation, and precipitation. There are eight wells that not proper for drinking and six wells that proper for drinking. Integration lab result, stratigraphic analysis, and depth aquifer show that groundwater that proper for drinking comes from confined aquifer while that not proper for drinking comes from unconfined aquifer.

Keywords: Facies, Gibbs, Hydrogeochemical, Piper

1. Introduction

ITERA is a first new state technology institute at Sumatera. ITERA from 2014 till 2018 has been lecturing for about five thousand students. This number will be increasing as long as new students acceptance every year. The additional of amounts students will trigger economic growth, especially at campus surroundings. To anticipate economic growth in the future, research is very important to be done to understand groundwater hydrogeochemical characteristic of the research area and its proper to all needs especially for drinking water purpose (Chang and Wang, 2010; Wen et al., 2005). It is based on that water is the most important element in human health.

Groundwater chemistry composition depends on hydrogeochemical process that groundwater pass. Groundwater chemical compositions are integration natural and anthropogenic factors such as precipitation, oxidation-reduction between groundwater and mineral aquifers, geological structures, cation exchange, mineral dissolution, water mixing, fertilizer leaching, biology process, and human activities. All these interactions result in variations of groundwater type (Yang et al., 2016). Thus, hydrogeochemical study what types of process that control groundwater hydrochemistry (Jeevanandam et al., 2007).

The objects of research including groundwater from dig well, that could be called as groundwater from unconfined aquifer and groundwater from drill well, that could be called as groundwater from confined aquifer. The research purposes are to interpret hydrogeochemical processes that control groundwater chemistry composition and groundwater assessment for drinking purpose. This research might be the first research that ever been done at this research area. Hopefully, this research result could help local government to make policies about development area in the future.

2. Geology and Hydrogeology

Research area is included within Tanjungkarang sheet regional geology map scale 1:250.000 (Mangga, S.A; Amirudin; Suwanti, T; Gafoer, 1993). The area is composed of Lampung Formation of Quaternary age (fig 1). The Lampung formation is composed of pumice tuff, rhyolitic tuff, tuff unified tuffit, tuffaceous claystone, and tuffaceous sandstone. From this, all lithologies are volcanic activities associated. Lampung Formation is deposited unconformity above of Andesite of Tertiary age. At above of Lampung Formation is deposited unconformity young volcanic deposits (lava andesite-basalt, breccia, and tuff). Geological structures that have been developed so little or still unidentified.

4. Result and Discussion

4.1 Hydrogeochemical Study

The groundwater observations had been conducted on 14 wells including nine dig wells and five drill wells (fig 2). All of nine dig wells having range water depth between 0,2 m-8,2 m which can be classified into unconfined aquifer (Table 1). While other five drill wells having well depth about 40 m and 80 m which can be classified into confined aquifer. It is also supported by previous research result that had been conducted at campus ITERA in 2017. The previous research was purposed to detect type of aquifer and its depth with Schlumberger configuration-geolectrical method. This research concluded, stratigraphically, the rocks from younger to oldest are siltstone, claystone, sandstone, and claystone with depth aquifer at more than 25 m (Setiawan et al., 2017). If observed from rocks ability to storage, to release water and also connected to its position to upper rock layer and lower rock layer, then can be concluded that sandstone is confined aquifer.

Piper diagram is one of the most effective graphic representation in the study of the groundwater quality, which helps to understand the groundwater geochemical characteristics (Yang et al., 2016). Based on plotting data of cation and anion into Piper diagram,

there are five hydrochemical facieses, they are Na-HCO₃-Cl (35,71%), Na-Cl (21,43%), Na-HCO₃ (21,43%), Na-SO₄-Cl (14,29%), and Ca-Mg-HCO₃ (7,14%) (fig 3).

4.2 Groundwater Assessment for Drinking Purpose

Groundwater assessment for drinking purpose at research area is according to Indonesia Healthy Ministry regulation standard No. 492/MENKES/PER/IV/2010 (Table 2). There are 18 parameters which be used as references for this research. The 18 parameters including 7 physics and 11 chemical data. Laboratory analysis result shows there are value differences which very contrast on some parameters. The parameters are colour, muddiness, TDS, Fe, pH, Mn, and NO₃ (Table 3 and Table 4). These contrast differences in value are evidently out of standard range which had been set. These are occurred on some samples were A1, A2, A3, A4, A5, A7, A8, A9 (Table 5).

X-Y plots are used to assess relative abundances of major cationic and anionic species present in different water environments (Pazand et al, 2018). There are four X-Y plots that will be used for analyzing the effect of rocks to groundwater. They are Na/Cl, Na+K / Ca+Mg, Ca+Mg / HCO₃, and Na+Cl / HCO₃+SO₄+Mg-Ca (meq/l) graphic.

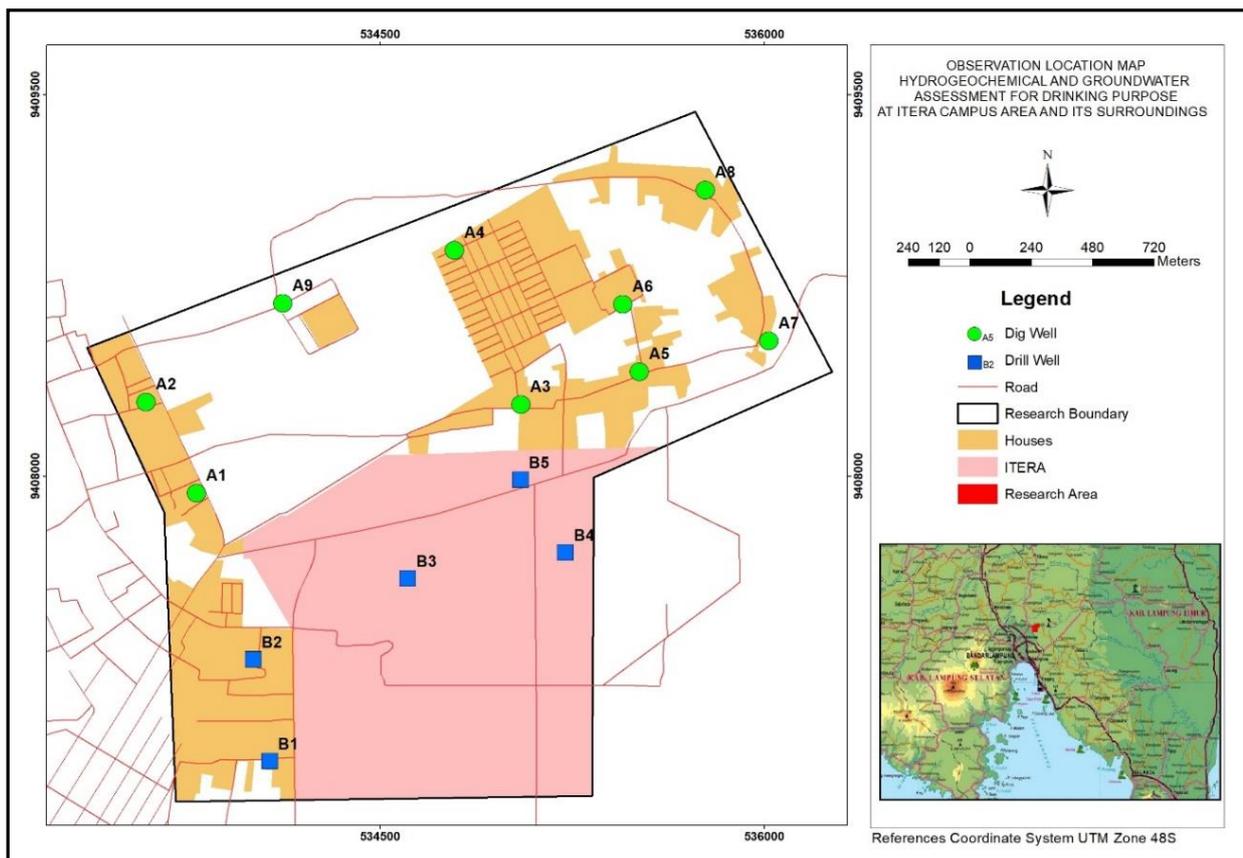


Fig. 2. The research area includes wells.

Table 1. The Aquifer types of research area based on integration with previous research.

Samples Code	Wells Type	Water Depth	Aquifer Type
A1	Dig	3	unconfined aquifer
A2	Dig	5,45	unconfined aquifer
A3	Dig	5,6	unconfined aquifer
A4	Dig	0,2	unconfined aquifer
A5	Dig	0,7	unconfined aquifer
A6	Dig	1	unconfined aquifer
A7	Dig	1,1	unconfined aquifer
A8	Dig	0,2	unconfined aquifer
A9	Dig	8,2	unconfined aquifer
B1	Drill	40	confined aquifer
B2	Drill	40	confined aquifer
B3	Drill	80	confined aquifer
B4	Drill	80	confined aquifer
B5	Drill	80	confined aquifer

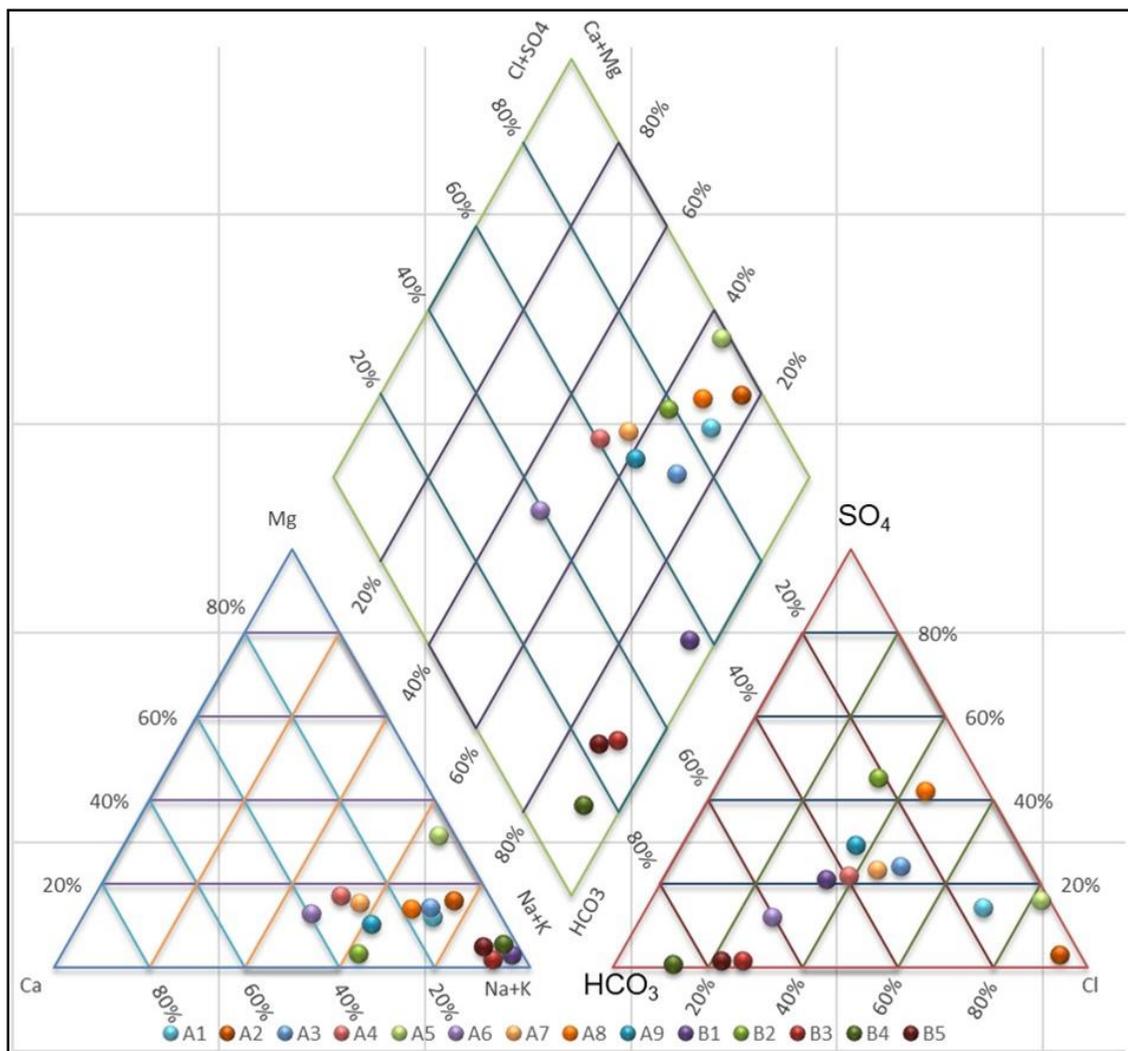


Fig. 3. Piper diagram of ionic compositions of groundwater in research area.

Na/Cl graphic is used to identify the mechanisms for acquiring salinity and saline intrusions in semi-arid regions (Yang et al, 2016). Na/Cl ratio is >1 , indicating that weathering of silicate rocks such as granodiorite, andesite, rhyolite and tuff was the primary process responsible for the release of Na^+ into the groundwater. Na/Cl ratio is <1 , the ion exchange and/or evaporation were dominant process resulting in the addition of Cl^- in the groundwater (Meybeck, 1987). Based on ion Na and Cl plotting into Na/Cl graphic (fig 4a), could be known that the ratio is >1 . This result shows that the ion Na comes from the weathering of silicate rocks.

The $(\text{Ca}^{2+} + \text{Mg}^{2+}) / \text{HCO}_3^-$ ratio is used to define the sources of Ca^{2+} and Mg^{2+} in groundwater (Fig 4b). If Ca^{2+} , Mg^{2+} and HCO_3^- in waters are derived from carbonate minerals, the ratio of $(\text{Ca}^{2+} + \text{Mg}^{2+}) / \text{HCO}_3^-$ should equal to 1 (Zhang et al, 2015). Fig 4b shows that the ratio was not equal to 1, so can be known that the source of Ca^{2+} and Mg^{2+} come from another source. Fig 4c shows that Na^+ and K^+ are relatively more abundant than Ca^{2+} and Mg^{2+} . It was associated with volcanic terrain and sourced from the weathering of K-feldspar and Plagioclase. Na+Cl / $\text{HCO}_3^- + \text{SO}_4 + \text{Mg} + \text{Ca}$ graphic used to identify the mechanisms for obtaining cation exchange and adsorption. If there are cation exchange and adsorption, the point is close to the 1:1 line (Pazand et al, 2018). Fig 4d shows $R^2 = 0.8126$ indicating there are different degrees of cation exchange adsorption in study area.

Gibbs plot shows as a function of the TDS that has the ability to provide information about the relative importance of the major natural mechanisms controlling groundwater chemistry and is extensively used to assess the functional sources of dissolved chemical constituents, such as precipitation dominance, rock dominance, and evaporation dominance (Pazand et al, 2018). Fig 5 shows that groundwater chemistry is mainly controlled by rock weathering and balance of evaporation-precipitation condition.

The chemical of groundwater including ion changing or chemistry reactions can be changed as time goes by. Next research needs to be done to know this changing in a certain period.

4.2 Groundwater Assessment for Drinking Purpose

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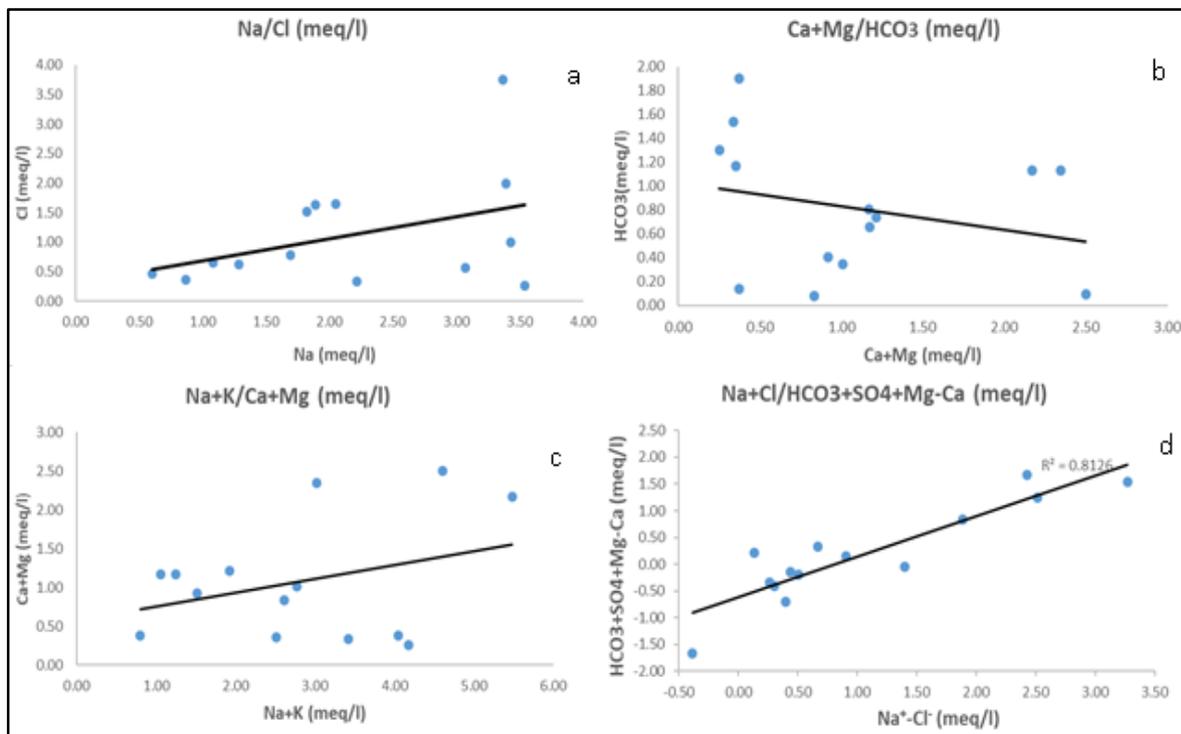


Fig. 4. Ions scatter diagram of groundwater in the study area.

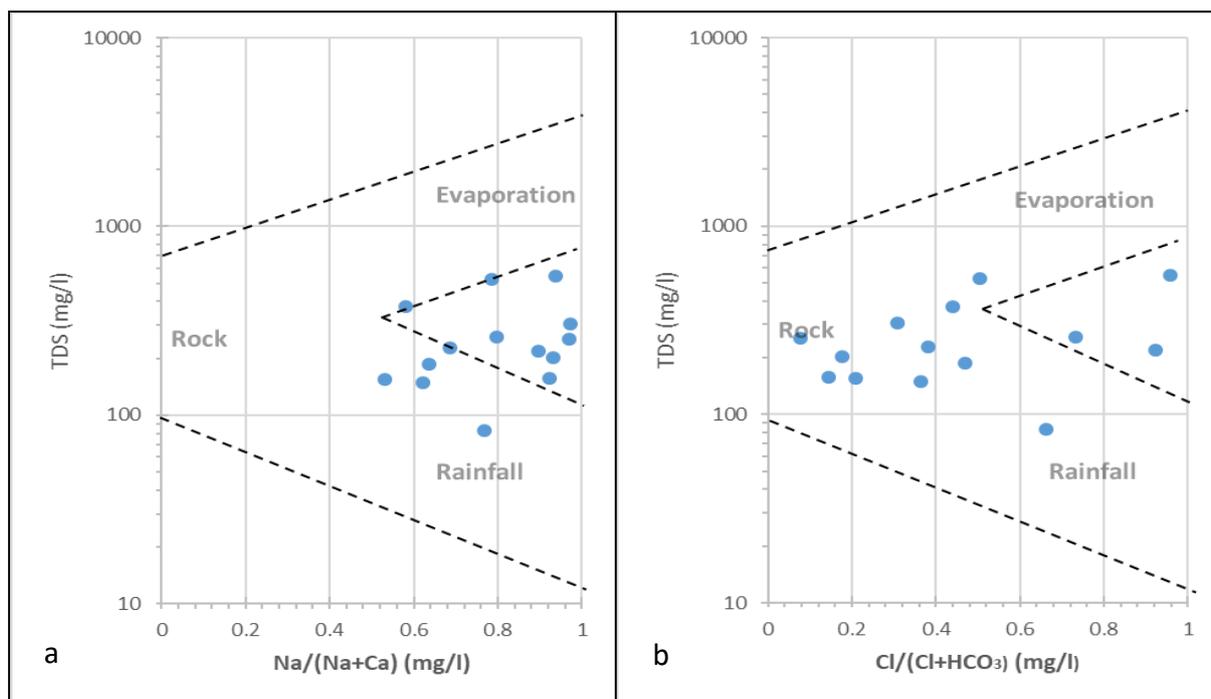


Fig. 5. Relationships between ion concentrations for Na^+ , Ca^{2+} , Cl^- , HCO_3^- with TDS.

Colour and Muddiness have straight correlation according to data. The samples which had very high value for colour are A4 (10 Pt. Co), and A8 (20 Pt. Co). The colour for A4 is yellow to red, it is affected by iron contamination and the colour for A8 is yellow to brownish it might be affected by iron mixing with organic matter. This result also occurred for muddiness parameter. Both samples also had high muddiness above standard, A4 (29,7 NTU) and A8 (65,4 NTU). Besides that, A7 (15,7 NTU) and A9 (7,66 NTU) also has high value but not higher than A4 and A8.

TDS depends mainly on the concentration of major ions such as HCO_3^- , SO_4^{2-} , Cl^- , Mg^{2+} , and Na^+ (Chang and Wang, 2010). The high value above standard for TDS parameter come from A3 (526 mg/L) and A5 (525 mg/L). These results are affected because of Cl^- and SO_4^{2-} ions that its source from tuff as volcanic deposits and/or anthropogenic contamination.

The rise in iron contamination in natural water sources is linked to various processes, which including oxidation-reduction reactions from weathering of iron rich minerals, microbiological activities, and anthropogenic iron contaminations (Sarkar and Shekhar, 2018). There are two samples that have high value of Fe, A4 (1,04 mg/L) and A8 (0,581 mg/L). These only about 14% of all samples. This percentage shows that this as a local phenomenon because of uncovering all area. This phenomenon shows the high value of Fe is effected by anthropogenic iron contaminations.

The presence of Mn as same as with Fe. Both presence due to either natural or anthropogenic sources (Corniello and Ducci, 2014). Natural sources come from weathering of minerals (pyroxenes, amphiboles, biotite, magnetite and in particular, olivine). While, anthropogenic sources come from wastewater discharge, dust and aerosols during

metallurgical processing, coal combustion, corrosion of water pumping infrastructure and transport of minerals or contamination associated with mining activities (Esteller et al, 2017). The main factors controlling the presence of these elements in water are pH, redox conditions and presence of organic or inorganic ligands (Corniello and Ducci, 2014). An acidic pH indicates that both ions are mobile, while a more neutral pH indicates that mobility is determined by redox conditions (Esteller et al, 2017). The Table 4 shows that A2 and A5 have direct correlation with pH value (acid) so could be concluded that Mn source comes from anthropogenic not by redox conditions.

Typical sources of nitrate in groundwater are mainly related to agricultural and domestic wastewater discharges (Andersen and Kristiansen, 1983). There were three samples which had higher result than standard, A2 (59,8 mg/L), A3 (55,7 mg/L), and A5 (114 mg/L). These covered about 21% of all samples. These results might be had direct correlation with research area that mostly was covered by agriculture about 60%.

The samples which contain pH out of standard are A1, A2, and A5. All of them contain pH under 6. A2 and A5 might be affected by dominance the presence of Mn and Fe. A1 might be affected by intake CO_2 from Atmosphere.

Table 5 shows that almost all samples especially which are taken from dig wells cannot be used for drinking water but samples which are taken from drill wells can be used for drinking purpose. This result also gives information that groundwater from unconfined aquifer is not good for drinking purpose but groundwater from confined aquifer is good for drinking purpose.

Table 2. The groundwater physics and chemistry standard for drinking purpose based on No. 492/MENKES/PER/IV/2010.

Parameters	Standard
Smell	no smell
Flavour	no flavour
Colour (Pt.Co)	15
Muddiness (NTU)	5
EC ($\mu\text{S}/\text{cm}$)	no information
TDS (mg/L)	500
Temp ($^{\circ}\text{C}$)	$\pm 3^{\circ}\text{C}$
Fe (mg/L)	0,3
F (mg/L)	1,5
pH	6,5-8,5
Mn (mg/L)	0,4
NO_3 (mg/L)	50
NO_2 (mg/L)	3
Cl (mg/L)	250
SO_4 (mg/L)	250
Na (mg/L)	200
CaCO_3 (mg/L CaCO_3)	500
Pb (mg/L)	0,01

Table 3. The physical data of all samples at research area.

Samples Code	Wells Type	Smell	Taste	Colour (Pt. Co)	Muddiness (NTU)	EC ($\mu\text{S}/\text{cm}$)	TDS (mg/L)	Temp ($^{\circ}\text{C}$)
A1	Dig	No	No	5	2,41	433	259	25,7
A2	Dig	No	No	5	0,07	365	219	25,8
A3	Dig	No	No	5	1,51	752	526	25,7
A4	Dig	Yes	Yes	10	29,7	248	149	25,7
A5	Dig	No	No	5	4,94	778	545	25,7
A6	Dig	No	No	5	2,48	222	155	25,7
A7	Dig	No	No	5	15,7	535	375	25,7
A8	Dig	Yes	Yes	20	65,4	138	83	24,8
A9	Dig	No	No	5	7,66	326	228	25,7
B1	Drill	No	No	5	0,79	435	304	25,9
B2	Drill	No	No	5	1,14	311	187	25,8
B3	Drill	No	No	5	0,81	339	203	25,8
B4	Drill	No	No	5	2,3	423	253	24,7
B5	Drill	No	No	5	0,84	261	157	25,8

Table 4. The chemistry data of all samples which taken at research area.

Samples Code	Wells Type	Fe (mg/L)	F (mg/L)	pH	Mn (mg/L)	NO ₃ (mg/L)	NO ₂ (mg/L)	Cl (mg/L)	SO ₄ (mg/L)	Na (mg/L)	CaCO ₃ (mg/L CaCO ₃)	Pb (mg/L)
A1	Dig	0,01	0,543	5,86	<0,2	45,2	0,3	57,7	15,9	43,5	50,5	< 0,001
A2	Dig	0,01	0,063	5,3	0,458	59,8	0,006	58,6	2,76	47,1	42	< 0,001
A3	Dig	0,045	0,273	7,1	<0,2	55,7	0,004	70,6	48	78	110	< 0,001
A4	Dig	1,04	0,123	6,43	<0,2	7,41	0,004	22,9	17,8	24,9	59	< 0,001
A5	Dig	0,01	0,255	5,54	1,11	114	0,665	133	35,7	77,4	126	< 0,001
A6	Dig	0,01	0,162	6,53	<0,2	6,85	0,004	12,9	7,89	20	59	< 0,001
A7	Dig	0,232	0,255	6,5	<0,2	22,7	0,117	53,9	38,8	41,9	118	< 0,001
A8	Dig	0,581	0,233	6,64	<0,2	5,45	0,004	16,6	21,5	13,9	19	< 0,001
A9	Dig	0,01	0,181	6,58	<0,2	25,9	0,073	27,8	30,3	38,9	61	< 0,001
B1	Drill	0,01	0,409	7,06	<0,2	3,88	0,004	35,5	29,8	78,8	12,5	< 0,001
B2	Drill	0,01	0,103	6,33	<0,2	16,9	0,004	21,9	41	29,6	46,3	< 0,001
B3	Drill	0,172	0,457	7,15	<0,2	2,93	0,004	19,9	1,93	70,7	17	< 0,001
B4	Drill	0,01	0,574	7,38	<0,2	2,29	0,391	9,58	1	81,4	19	< 0,001
B5	Drill	0,01	0,291	6,83	<0,2	2,09	0,004	11,9	1,41	51,0	18	< 0,001

Table 5. List of samples out of standard

Parameters	Min	Max	Samples out of standard
Colour (Pt.Co)	5	20	A4, A8
Muddiness (NTU)	0,07	65,4	A4, A7, A8, A9
TDS (mg/L)	83	545	A3, A5
Fe (mg/L)	0,01	1,04	A4, A8
pH	5,3	7,38	A1, A2, A5
Mn (mg/L)	0,2	1,11	A2, A5
NO ₃ (mg/L)	2,09	114	A2, A3, A5

5. Conclusions

From research at ITERA campus area and its surroundings then can be concluded as follow:

- Groundwater consist of five groundwater facieses, they are Facies Na-HCO₃-Cl (35,71%), Facies Na-Cl (21,43%), Facies Na-HCO₃ (21,43%), Facies Na-SO₄-Cl (14,29%), and Facies Ca-Mg-HCO₃ (7,14%).
- X-Y plots show that Na⁺ and K⁺ at research area come from weathering of silicate minerals. Gibbs plot shows that there is another factor that controls groundwater chemistry in addition to effect from rocks weathering. The factor is a balance of evaporation-precipitation condition.
- Groundwater which comes from unconfined aquifer (dig wells) is not proper for drinking purpose because it has been polluted by effect of human activities. Besides that, well condition without roof or cap makes groundwater be contaminated by precipitation easily. Precipitation can makes pH groundwater decreasing.
- Groundwater which comes from confined aquifer **with depth ≥ 40 m and closed well condition (drill well)** is a good source for drinking purpose.

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